Improvement of vertical profiles of raindrop size distribution from micro rain radar using 2D video



disdrometer measurements



Elisa Adirosi (1,2), Luca Baldini (1), Nicoletta Roberto (1), Patrick Gatlin (3), Ali Tokay (4,5), Daniele Casella (1)

MRR Processing

consider the two adjacent spectra. Starting from the

peak of the j-th spectrum the lower (or upper) limit of

the j-th spectrum is the first bin between the peak of

the j-th spectrum and the peak of the (j-1)-th (or

(j+1)-th) spectrum with a value of spectral reflectivity

lower than the noise level, or, if it does not exist, with

the bin corresponding to the minimum values of the

(1) Institute of Atmospheric Sciences and Climate (CNR), Rome, Italy (2) Department of Civil, Building and Environmental Engineering, Sapienza University of Rome, Italy, (3) NASA Marshall Space Flight Center, Huntsville, Alabama, (4) Joint Center for Earth Systems Technology, University of Maryland Baltimore County, Baltimore, Maryland, (5) NASA Goddard Space Flight Center, Greenbelt, Maryland

Rationale

- ☐ A number of RSD datasets collected with disdrometers (such as 2D video disdrometer, **2DVD**) and profilers (such as Kband vertically pointing micro rain radar, MRR) are available in several climate regions from GPM GV campaigns, but few related studies have utilized the MRR. This is mainly due to the questionable validity of standard MRR products in heavy precipitation or convection.: Vertical winds, attenuation correction, Doppler spectra aliasing, and range-Doppler ambiguity can limit the performance of MRR in these conditions.
- A precarious assumption used in the MRR standard processing to retrieve the RSD is that the vertical wind is negligible. This assumption is not very appropriate during convection, where the presence of up/down draft can shift the measured velocity spectrum leading to an inaccurate retrieval of the RSD and thus bias integral parameters..

MRR raw spectra

"("Raw spectra" file

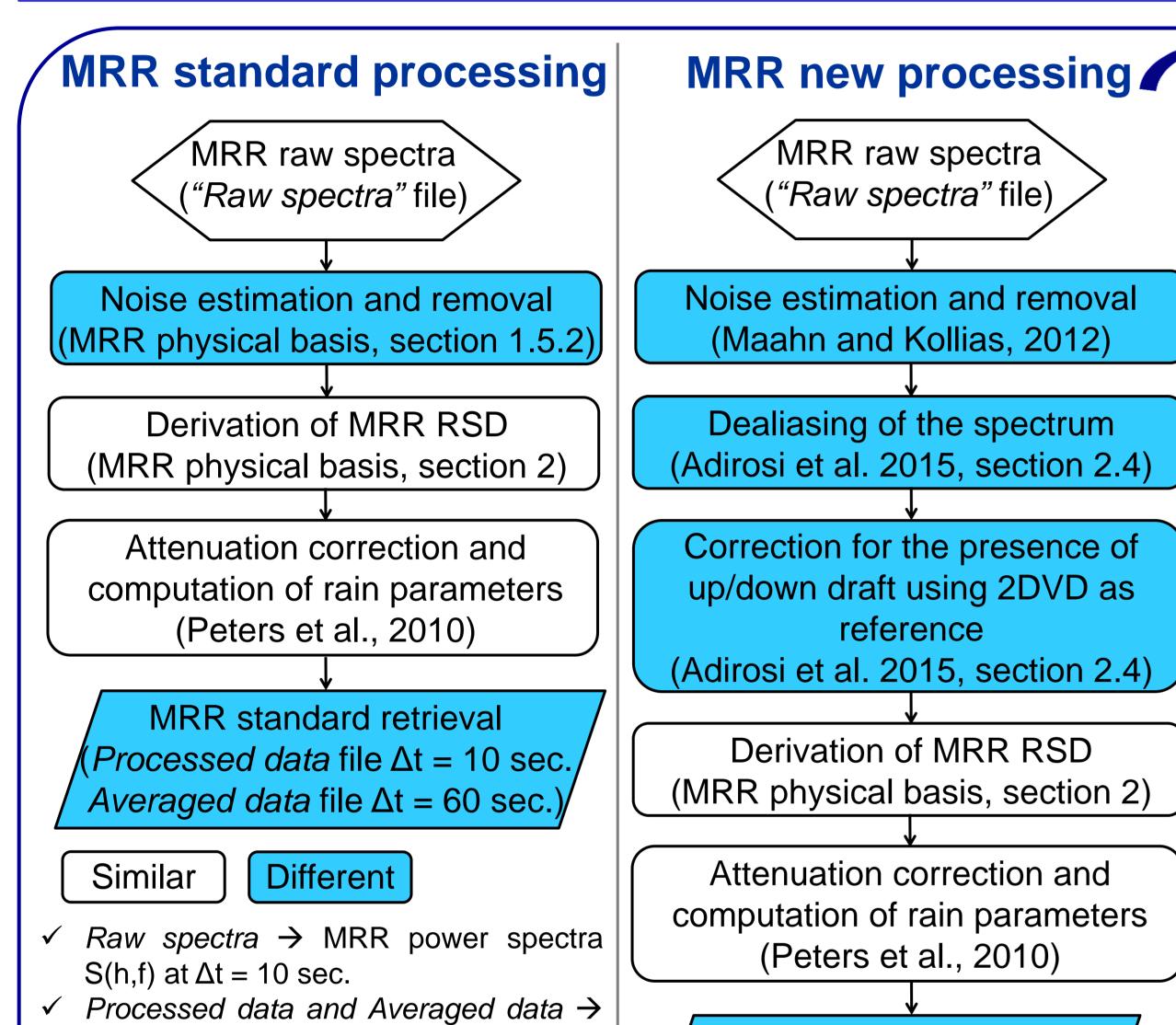
reference

(Peters et al., 2010)

MRR new retrieval

Objectives

- Datasets of MRR spectra can be properly reprocessed exploiting the synergy with coincident 2DVD measurements to reduce uncertainty of MRR profiles estimates and to increase the reliability of MRR data.
- > In particular, MRR Doppler spectra are reprocessed, exploiting the 2DVD measurements at ground to estimate the effects of vertical winds at 105 m (the most reliable MRR lower height), to provide a better estimation of vertical profiles of RSD.
- > MRR and 2DVD installed close each other at a site in the historic center of Rome (Italy) during the Special Observation Period 1 of the HyMeX (Hydrological Cycle in the Mediterranean Experiment) project, (5 September - 6 November 2012) are used to evaluate the proposed method.



corresponding integral

event averaged RSD

comparison

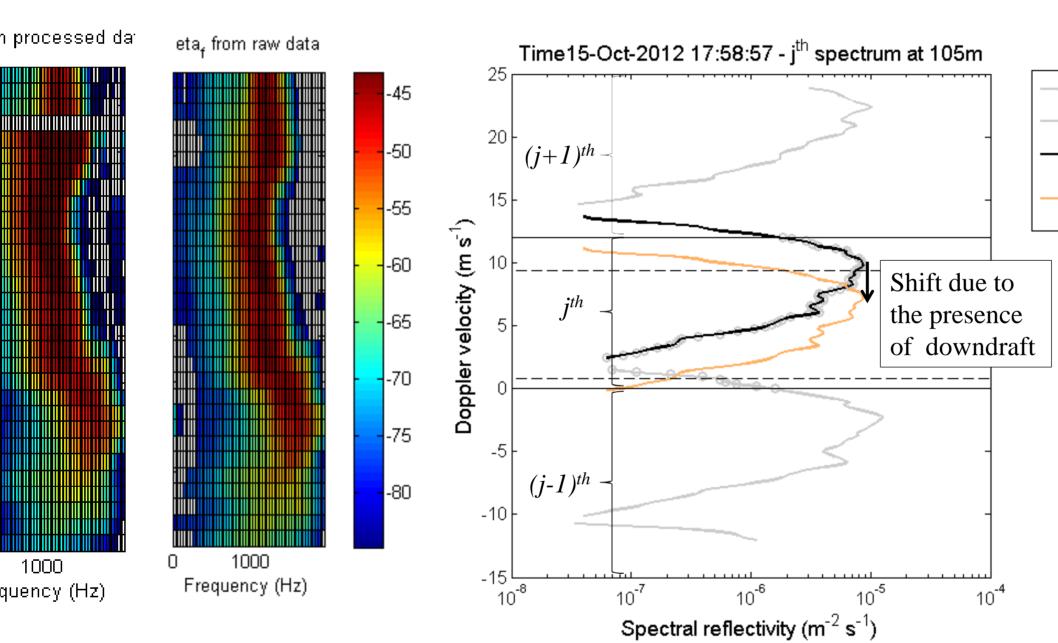
RSD and

Mean RSD day 1031

variables.

Step 1: noise removal

Example of reflectivity spectra Processed from standard data file (15 October 2012 at 18:01:27 UTC) (left) and those obtained applying the noise removal procedure by Maahn and Kollias (2012) to raw spectra (right).



spectral reflectivity.

Step 3: correction for up/down daft Step 2: dealiasing To unfold the spectrum of the j-th range gate we

The characteristic fall velocity of 2DVD

The characteristic fall velocity of 2DVD
$$v_{c,2DVD}(D) = \frac{\int_{D_{min}}^{D_{max}} v_t (D) N(D) \, \sigma_b(D) \, dD}{\int_{D_{min}}^{D_{max}} N(D) \, \sigma_b(D) \, dD}$$

Is matched to the one obtained from the MRR spectra at 105 m

$$v_{c,MRR} = \frac{\int \eta_v(v) \, v \, dv}{\int \eta_v(v) \, dv}$$

N(D) = 2DVD measured RSD, $v_t(D) = terminal fall$ velocity, $\sigma_b(D)$ = backscattering cross section area computed at 24.243 GHz, and $\eta_{\nu}(v)$ = dealiased Doppler spectrum of MRR.

The velocity difference at 105 m has been used to correct all the range gates. Assumptions:

- > the vertical change in RSD within the lowest 100 m is negligible
- > the vertical motions are near zero at the height of the 2DVD measurements (at ground)
- > the up/down draft does not vary significantly within the lowest 1 km AGL (e.g. Giangrande et al. 2013).

Step 4: RSD retrieval

At all the different elevations, the estimated been normalizing the spectral reflectivity density with respect to drop diameter $\eta_D(i,j) = \eta_v(i,j) 6.18 \exp(-0.6 D) \delta(h)$

by
$$\sigma_b(D)$$

$$N(D,i) = (\eta(D,i))/(\sigma(D))$$

Step 5: attenuation correction

technique attenuation correction by Peters et al. (2010) was applied to processed raw spectra. path integrated Spectra attenuation greater than 10 dB, were discarded.

MRR new processing (NEW@105), and

Validation of MRR new processing

 ○ AVE@105 → Reflectivity or RSD from MRR standard processing in the Average data file at 105 m ○ NEW@105 → Reflectivity or RSD from the MRR new processing at 105 m

step 2

step 3

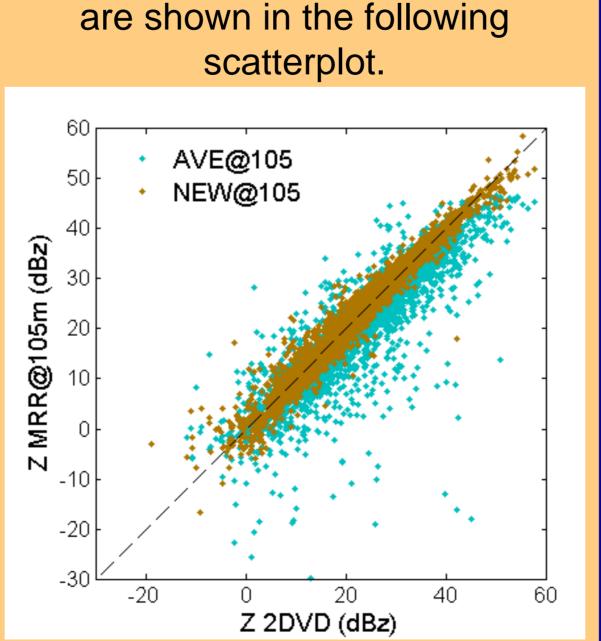
step 5

---AVE@105 ---2DVD data Reflectivity comparison C/S classification Z_{2DVD} vs Z_{NEW@105m} Z_{2DVD} vs Z_{AVE@105m} minutes ---AVE@105 ---2DVD data 0.04(0.17) 0.965

• Normalised standard error: $NSE = \frac{\sqrt{(x-y)^2}}{-}$ • Normalised Bias: $NB = \frac{\sum y}{\sum x} - 1$ (NB < 0 means underestimation of MRR with respect to the 2DVD)

• Correlation coefficient: cc =

The Stratiform/Convective (C/S) algorithm proposed by Bringi et al. (2009) has been applied to the 2DVD data ...



The reflectivities of all the events

The proposed processing allows a **better agreement** between MRR and the 2DVD.

Box plot of Z and D_0 at the different heights during stratiform rain (on the left), and convective rain (on the right).

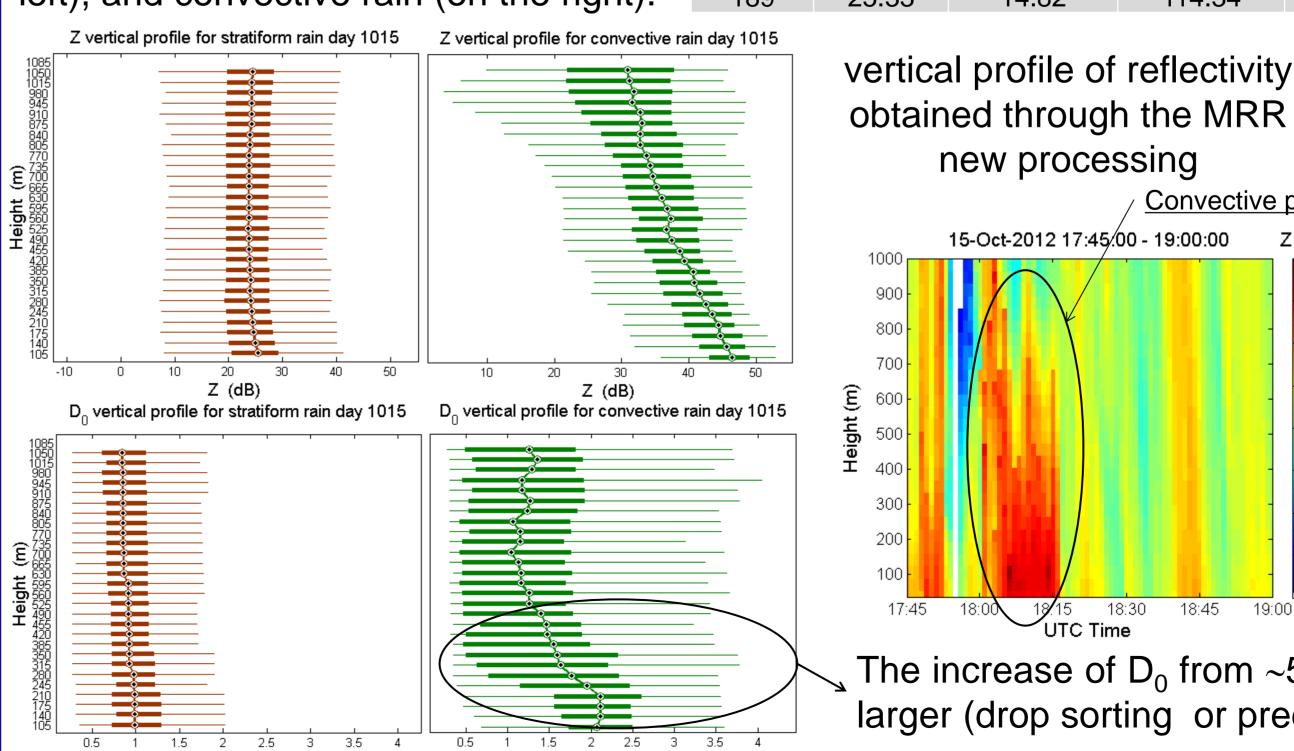
jth spectrum

(j-1)th and (j+1)th spectra

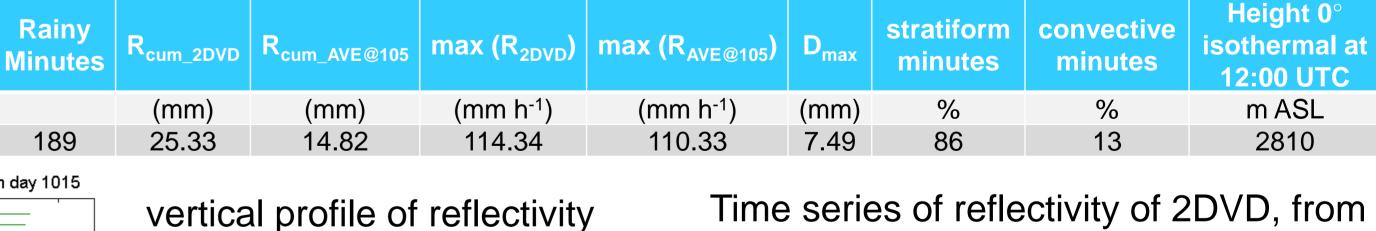
th dealiased spectrum

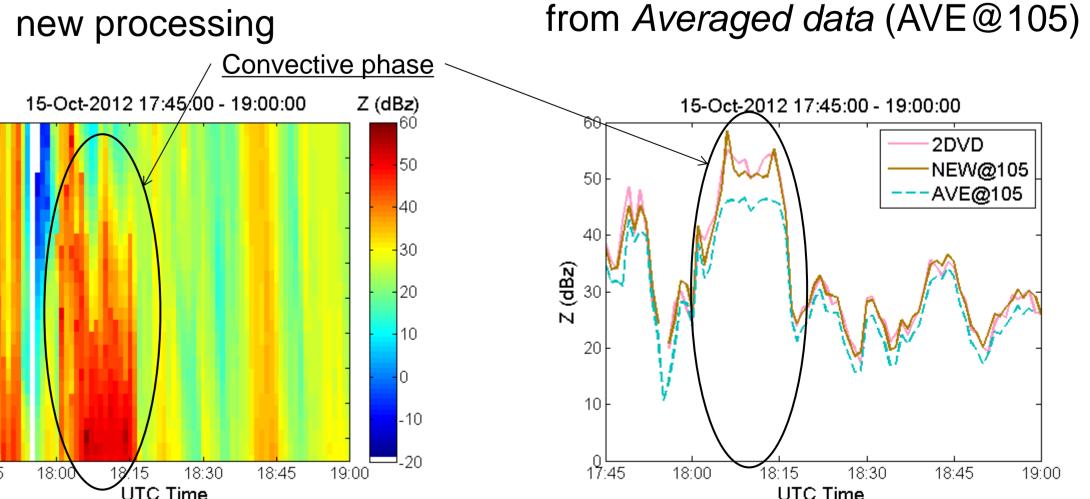
+ vertical wind correction

-jth dealiased spectrum



Case study: 15 October 2015





The increase of D₀ from ~500m AGL to the ground means drops become larger (drop sorting or predominance of coalescence).

References: Adirosi, E., Baldini, L., Roberto, N., Gatlin, P, Tokay, A., Atmos. Res., in press (2015); Giangrande, S. E., Collis S., Straka J., Protat A., Williams C., Krueger S., J. Appl. Meteor. Climatol., 52, 2278–2295 (2013); Maahn, M., Kollias, P., Atmos. Meas. Tech., 5, 2661-2673 (2012); METEK, 2012., MRR Physical Basics, (Update of 13 March 2012) Elmshorn, 20 pp., Metek mbH; Peters, G., Fischer, B., Münster, H., Clemens, M., J. Atmos. Ocean. Technol., 44, 829-842 (2010).